NEAR REAL TIME FLOOD FORECASTING IN THE TRANSBOUNDARY CHENAB RIVER USING GLOBAL SATELLITE MAPPING OF PRECIPITATION

Sikandar Ali¹, *, Muhammad Jehanzeb Masud Cheema², Allah Bakhsh¹, Tasneem Khaliq³

¹Department of Irrigation and Drainage, University of Agriculture, Faisalabad, Pakistan; ²Faculty of Agricultural Engineering and Technology, Pir Mehr Ali Shah Arid Agriculture University, Rawalpindi, Pakistan; ³Department of Agronomy, University of Agriculture, Faisalabad, Pakistan

*Corresponding author’s e-mail: sikandar.ali@uaf.edu.pk

INTRODUCTION

Floods are the most ruinous natural catastrophes among all the natural disasters (Stefanidis and Stathis, 2013) and cause various socio-economic skirmishes (Ghimire et al., 2015). Floods affected more than 8 million people in the 20th century worldwide (Haynes et al., 2009), disturbed livelihoods of about 520 million people and claimed 25000 lives annually across the world (Simonovic, 2009). In Pakistan, the super flood destroyed about 5.189 million acres of agricultural land in 2010 and caused huge economic loss to the country (OA, 2010). The Chenab river flood affected 2.415 million acres of standing crops in 2014 mainly covering cotton, rice, sugarcane and vegetables with direct damage to a large number of water channels and wells (FFC, 2017). Due to 2010 flood, the major crops like rice, wheat, cotton and sugarcane led to a declined growth of 4% and the overall agriculture sector recorded modest growth of 2.1% against the target of 3.8% (Looney, 2012) which caused huge economic dent and significant recovery cost. The research conducted by Multihazard Mitigation Council and United Nations Development Programme reveals that the money spent on the development of early warning system is inexpensive than the recovery cost (MMC, 2005; UNDP, 2012). Improvement in flood forecasting mechanism is one of the most effective forms of non-structural approaches which could offer a vital role in flood preparedness and disaster management (Yang et al., 2015).

Real time flood forecasting is a major challenge to the researchers worldwide due to prevailing conditions over the high altitude complex river basins which are influenced by both the sensitivity of hydrologic models and climatic information of the area (FFC, 2000; Groisman et al., 2004; Tariq and Giesen, 2012). The flood forecasting in the transboundary river, like the Chenab river, is more complicated due to delay of data transmissions among the users (Biemans et al., 2009) especially in the snow dominant catchments (Herold et al., 2016). The HEC-HMS model can simulate the river flows based on any individual precipitation or continuous time period. Several hydrologic models have been applied in the river basins for event-based and continuous streamflow simulation across the globe (Anderson et al., 2002; Cydzik and Hogue, 2009; De Silva et al., 2014; Teng et al., 2018; Ouedraogo et al., 2018; Darianeet al., 2020). The HEC_HMS model can also be applied for flood forecasting, flood frequency, system planning and early warning in the complex river basins (HEC, 2015). The SMA and temperature index are the suitable methods for loss
estimation and snowmelt runoff modeling, respectively, in HEC-HMS model (Sintayehu, 2015; García et al., 2008; Wallner et al., 2012).

The flood forecasting cannot be performed without estimation of relatively highly accurate precipitation estimates (Yang et al., 2015) which is the key parameter among all the meteorological parameters (Tapiador et al., 2012). Due to sparse network of raingauge stations and absence in high-altitude areas of the world (Immerzeel et al., 2012), there is a need to use timely available satellite precipitation for flood forecasting. There are precipitation estimates derived from the constellation of passive microwave radiometers working on low-earth orbiting satellites and often merging data from geostationary infrared datasets for minimizing temporal uncertainties (Huffman et al., 2007). The GSMaP is the project established by the Japan Science and Technology Agency and Japan Aerospace Exploration Agency to produce global precipitation at a spatial resolution of 0.25° and 0.1° using merged microwave radiometer and precipitation radar algorithms. The project aims to develop an advanced microwave algorithm which is based on a deterministic rain retrieval algorithm (Aonashi et al., 2009) based on physical precipitation models consisting of particle size distribution and melting layers. (Kubota et al., 2007).

Flood forecasting of the transboundary Chenab river is a major challenge due to lack of hydrometeorological data of its catchment. In case of Chenab river, Pakistan is at downstream and being an agricultural country timely information of floods is very crucial for agricultural water management. In this study, an innovative and comprehensive approach is discussed to forecast floods in the Chenab river using satellite-based precipitation estimates. The three high spatiotemporal GSMaP precipitation estimates were integrated into the HEC-HMS model for flood forecasting of the Chenab river. The event-based and continuous simulation was performed using global temperature, snow depth, snow density, landuse and soil datasets along with the GSMaP precipitation estimates. The accuracy of flood forecasting was assessed by statistical comparison between the observed and simulated flows at the outlet (Marala barrage).

**MATERIALS AND METHODS**

**Study Area:** The Chenab river catchment lies in the geographic coordinate range of 74° to 77.83° E longitude and 32.1° to 34.1° N latitude with an elevation range of 224m to 7085m above mean sea level as shown in figure 1. The Chenab river originates from the Lahul-Spiti district of India, enters Jammu and Kashmir in Kishtwar district and then flows towards Pakistan in Sialkot district by covering a total length of about 590 km from its origin point to the outlet. For modeling purpose, the catchment of Chenab river is divided into 20 subbasins for hydrological modeling as shown in figure 1. The high-altitude of the catchment receives heavy snow and extreme precipitation in winter and summer season, respectively. The monsoon season is overlapped with the summer season which can cause extreme floods in the Chenab river due to heavy precipitation and snowmelt at the same time in its catchment.

**Datasets used in the study:** In this study, precipitation estimates with different temporal resolutions and latency periods were incorporated in the HEC-HMS model for simulating the daily Chenab river flow. The GSMaP is the highly precise and high spatiotemporal resolution precipitation estimate consisting of extensive satellite information from both infrared and passive microwave radiometer (Okamoto et al., 2005; Kubota et al., 2007; Aonashi et al., 2009; Ushio et al., 2009). The first precipitation estimate GSMaP_NRT was used in the study with a spatial resolution of 0.1° on hourly timespans with a very low latency period which is the core cause for its use in monitoring the hydrological response in the river basins. The GSMaP_NRT precipitation estimate has various input data sources (Ebert et al., 2007) and offering its availability to the public since 2002. The GSMaP_NOW half-hourly precipitation estimate is also used with the same spatial resolution extrapolated by 0.5-hour using cloud motion vector and updated by every 30 minutes for water resources and flood management studies (Tsujimoto et al., 2014). The GSMaP_RNC precipitation forecast provides data for the next 12 hours (OTSUOKA et al., 2019) which could be used effectively for flood forecasting. The GSMaP_RNC is formulated by the extrapolation of GSMaP_NRT using the cloud motion vector patterns. The latency period of GSMaP_NRT and GSMaP_NOW precipitation product is 4
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The soil layers give wetting and recovery cycles for simulation of the watersheds which requires the initial soil moisture conditions of the subbasins (Alexakis et al., 2017; Bhuiyan et al., 2017). As most of the watershed of the Chenab river lies in Indian control as shown in figure 1, so alternate sources of information were incorporated for the catchment. The initial values for canopy storage, surface storage and tension storage were fixed by keeping in view the range mentioned in HEC-HMS technical reference manual (Schaffenberg and Fleming, 2006). The daily MODIS (MOD09GA) imagery and landuse of the European Space Agency (ESA) were also used in order to keep the initial values within the limits. The values of maximum infiltration, imperviousness, soil storage and soil percolation were extracted from the Food and Agriculture Organization (FAO) soil database, ESA landuse and daily MOD09GA product. The Soil Conservation Services (SCS) unit hydrograph, lag time and constant monthly methods were used for estimation of direct runoff, routing and baseflow, respectively.

**Temperature index method:** The temperature index method was used in HEC-HMS model for simulating the snowmelt runoff (Fazel et al., 2013; Razmkhah, 2016; Dariane et al., 2020), which defines the subbasin precipitation as either liquid water or frozen snow by defining a temperature threshold. The melted water acts as hyetograph for the corresponding subbasin and thereby contributes to the streamflow. The temperature index method represents cold energy stored in the snowpack (SWE) with past conditions and all the other factors for estimation of melting for each degree above freezing temperature. Every subbasin is divided into different elevation bands by using Advanced Spaceborne Thermal Emission and Reflection Radiometer Digital Elevation Model.

For calibration of the model, event-based calibration was carried out for 4 events occurring between 2017 and 2018 using SMA and temperature index methods. Initially, the automatic and manual calibration was applied for fixation of all the SMA and temperature index parameters. Afterward, the continuous snowmelt runoff simulation was carried out for the complete timespan ranging from 2017 to 2018. The simulated and observed flows were compared for accuracy assessment using the statistical approaches like Coefficient of Determination ($R^2$), Pearson r and NSE.

**RESULTS**

The first event was simulated from 29-01-2017 to 14-02-2017 using all the mentioned GSMaP precipitation estimates. The results in figure 3, supported by Table 1, show that there is a very poor agreement between the observed and simulated flow for GSMaP_RNC precipitation product. Whereas, the other 2 precipitation estimates show fair results with underestimation at peak observed flow time. The higher values of $R^2$, Pearson r and NSE for GSMaP_NOW based

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and 1 hour, respectively, (Kubota et al., 2017; Tang et al., 2017) while GSMaP_RNC forecasts precipitation 12 hours in advance.

The daily Climate Prediction Center (CPC) temperature dataset with a spatial resolution of 0.5° was incorporated in the study for snowmelt runoff modeling in HEC-HMS model. The CPC temperature dataset is formulated by more than 30,000 gauge reports obtained from multiple National and International organizations of the world (Xie et al., 2007). In addition, Moderate Resolution Imaging Spectroradiometer (MODIS) temperature information (MOD11A1) was also blended with the CPC temperature product at a spatial resolution of 0.5° for snowmelt runoff modeling. The European Centre for Medium-Range Weather Forecasts (ECMWF) Re-Analysis 5 (ERA5) was used (Hersbach and Dee, 2016) to estimate the snow properties over the river catchment and SWE was estimated using the equation (Sturm et al., 2010):

$$SWE = SD \frac{\rho_s}{\rho_w}$$

where, SD is snow depth, $\rho_s$ is the density of snow and $\rho_w$ is the density of water.

**Soil moisture accounting method:** The SMA method was used in the HEC-HMS model for retention by vegetative cover and simulation of water movement to the groundwater through soil layers (Feldman, 2000). The SMA algorithm describes the watershed in the series of storage layers as shown in figure 2.

![Figure 2. Description of Soil Moisture Accounting method (Feldman, 2000).](image-url)
flow simulation also indicate fair simulation results as compared to GSMaP_NRT based simulation results.

The second event was simulated for the time period ranging from 24-06-2017 to 17-08-2017, the results in figure 4 and Table 1 show a reliable agreement for GSMaP_NOW based simulation as compared to the other 2 simulations. The higher values of NSE for GSMaP_NOW based simulated flow also indicate reliable results. The GSMaP_NRT based simulation shows slightly fair results with underestimated simulated flow over most of the event time while poor results were obtained for GSMaP_RNC based simulation.

The third event was simulated from 26-06-2018 to 07-07-2018, figure 5 shows a very poor simulation response for GSMaP_RNC precipitation estimate, while the simulated flows of the other 2 precipitation products are following the pattern of observed flow except at peak flow time. The higher NSE value (0.49) of GSMaP_NRT based simulated flow is indicating fair results as compared to the GSMaP_NOW based simulation which is more deviating from the observed flow with NSE value of 0.37.

The fourth event was simulated from 04-08-2018 to 28-08-2018, figure 6 shows that there is good agreement of GSMaP_NOW based simulated flow with the observed flow as compared to the other 2 flow simulations. The GSMaP_NOW based simulated flow is under and overestimating the observed flow with lower values of R², Pearson r and NSE. While, the result of GSMaP_RNC based simulated flow is very poor with high underestimation revealing lower values of R², Pearson r and NSE. All the four events were used for calibrating the model and the automatic calibration performed well which provided information about the catchment behavior.
The calibrated model was further applied for long term continuous snowmelt modeling from 2017 to 2018 for all the mentioned GSMaP precipitation estimates. The results are presented in figure 7 (a-f), figures a, b and c are the time series plots for GSMaP_NOW, GSMaP_NRT and GSMaP_RNC based simulated flows, respectively. The figure 7 (d, e and f) shows statistical comparison, supported by Table 2, for GSMaP_NOW, GSMaP_NRT and GSMaP_RNC simulation, respectively. The GSMaP_NRT based simulated flow shows good agreement with the observed flow and it is also tested over the Indus basin by Sugiura et al. (2016). The GSMaP_NOW based simulated flow shows a slightly fair agreement with the observed flow for most of the timespan except for the peak flow time. In the summer of 2017, GSMaP_NOW simulation shows 1 peak while observed flow has three high magnitude flows and there is also variation in temporal occurrence of peak flow for summer season in 2018 as well. The GSMaP_RNC based simulated flow shows a very poor agreement with the observed flow with high over and underestimation for the entire time period of the simulation. The apparent reason for overestimation in precipitation magnitudes of GSMaP_RNC in tropics is due to overestimation of precipitation magnitudes by GSMaP_NRT (Kotsuki et al., 2019).

From figure 7 (b, c, e and f) the overestimation and underestimation of GSMaP_NRT and GSMaP_RNC based
simulated flows can also be observed with higher variations in the latter. The statistics in figure 7 (d-f) and Table 2 reveal that $R^2$ and Pearson r values are higher for all the simulated flows indicating reasonable reliability of GSMaP precipitation estimates for time series simulation. The GSMaP_NOW based simulated flow has the highest NSE which indicates fair comparison for most of the time periods with small variations persistently occurring in the entire timespan as can be seen in figure 7 (d). The variations in GSMaP_NRT based simulated flow can be seen for less number of days in figure 7 (e and f), with lower but reliable NSE value. The $R^2$ and Pearson r values for GSMaP_NRT based simulation are also reliable indicating a good precipitation source. The statistics of GSMaP_RNC based simulated flow indicates a very poor value of NSE, thereby indicating a poor precipitation forecast.

**Table 2. Statistical comparison of GSMaP products based simulated flows with the observed flow from 2017 to 2018.**

<table>
<thead>
<tr>
<th>Function</th>
<th>GSMaP_NOW</th>
<th>GSMaP_NRT</th>
<th>GSMaP_RNC</th>
</tr>
</thead>
<tbody>
<tr>
<td>$R^2$</td>
<td>0.86</td>
<td>0.80</td>
<td>0.71</td>
</tr>
<tr>
<td>Pearson r</td>
<td>0.85</td>
<td>0.77</td>
<td>0.65</td>
</tr>
<tr>
<td>NSE</td>
<td>0.73</td>
<td>0.58</td>
<td>0.35</td>
</tr>
</tbody>
</table>

**DISCUSSION**

The automatic calibration is an iterative approach used in the HEC-HMS model to reduce the uncertainty by optimizing the objective function which ultimately would reflect a good agreement between simulated and observed flows (Schaffenberg, 2013; Sintayehu, 2015; Majidi and Shahedi, 2012). The high spatiotemporal GSMaP_NRT, GSMaP_NOW and GSMaP_RNC precipitation products were used as input to the HEC-HMS model which are being applied worldwide (Aonashi et al., 2009; Ushio et al., 2009; Nakayama, 2015). For snow-related estimations, ERA-5 has been used by the researchers for cryosphere and flood studies across the globe (Terzago et al., 2019; Tei et al., 2019; Hersbach and Dee, 2016). The CPC and MODIS based temperature datasets were used in this research which are being used by the researchers due to their good results across the globe (Xu et al., 2013; Mallakpour and Villarini, 2015; Corbari and Mancini, 2014; Kwak and Park, 2019). The SMA method is also suitable for event based modeling (Bhuiyan et al., 2017) and long-time continuous hydrological modeling over the complex watersheds (Gyawali and Watkins, 2012; Singh and Jain, 2015). The SCS unit hydrograph method for estimation of direct runoff and lag time approach for routing was used in this study as recommended by the researchers worldwide (Verma et al., 2010; Majidi and Shahedi, 2012; Olivera, 2001; Razi et al., 2010). The temperature index method for snowmelt runoff modeling is a suitable approach in HEC-HMS model (Verdhen et al., 2013) and reliable results are observed by the researchers worldwide (Fazel et al., 2013; Azmat et al., 2017).

The GSMaP_NRT and GSMaP_NOW based simulations show reliable agreement with the observed flow for almost all of the time periods and tested by the researchers worldwide (Okamoto et al., 2005; Kubota et al., 2007). There are some uncertainties in simulations which can be minimized by local calibration of the precipitation estimates (Cheema and Bastiaanssen, 2012; Mashingia et al., 2014; Shiraishi et al., 2009) as self-corrected GSMaP_NRT has been applied over the Indus basin by Sugiuara et al. (2016). The GSMaP_RNC based simulated flow shows poor results with high over and underestimation as its overestimation in precipitation is proved in tropics by Kotsuki et al. (2019). The higher uncertainties are observed for GSMAP_NOW based simulation in this study as higher uncertainties are also observed for GSMaP_NOW precipitation by Kubota et al. (2017) which may lead to less accurate results due to decrease in microwave radiometer coverages with 0.5-hour latency. The variations in the peak magnitude and time of observed and simulated flow is also directly linked to the storage and release due to Indian dams on Chenab river. The accuracy of the simulation can be improved by timely adjustment of storage and release of water from the dams into the HEC-HMS model.

**Conclusion:** The purpose of this study was to develop a flood forecasting system for the transboundary Chenab river by integrating high spatiotemporal GSMaP precipitation estimates with the HEC-HMS model. The results indicated a reliable agreement between the observed and simulated flows for GSMaP_NRT and GSMaP_NOW precipitation estimates for time series simulation, while very poor simulation results are obtained by using GSMaP_RNC precipitation estimates. Due to their high spatiotemporal characteristics and low latency time, the GSMaP_NRT and GSMaP_NOW precipitation estimates can be used in hydrological models for real time flood forecasting. The real time flood forecasting would help to forecast the magnitude and volume of water which would be helpful in flood management in order to minimize the losses.

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